Lecture 7: Carbonate Geochemistry Continued

I. Review

A. Activity-concentration relationships

a. Equilibrium constants (K_{eq}) change as a function of temperature, resulting in increases or decreases in the ratio of products to reactants: \sim

$$
\ln(K_{T2}) - \ln(K_{T1}) = \frac{\Delta H_{rxn}^{\circ}}{R} \left(\frac{1}{T_1} - \frac{1}{T_2}\right)
$$

b. The saturation index (SI) describes whether a solution is undersaturated ($Q \leq K_{eq}$ and $SI \leq 0$) or oversaturated $(Q > K_{eq}$ and $SI > 0$) with respect to mineral:

$$
SI = log(Q/K_{eq})
$$

If the solution is undersaturated, mineral dissolution occurs; if the solution is oversaturated, mineral precipitation occurs.

B. The Bicarbonate Buffering System

We can explore bicarbonate buffering in the environment as an example of acid-base reactions.

 $CO₂$ gas dissolves in water until equilibrium is reached (or vice versa...e.g. when a soda can degasses open opening):

$$
CO_2 + H_2O \rightarrow H_2CO_3^*
$$

 H_2CO_3 ^{*} represents all forms of CO₂ dissolved in water $(H_2CO_3^* = H_2CO_3 + HCO_3^- + CO_3^2)$ We can use a series of equations to determine the concentration of H^+ as carbonic acid deprotonates.

The species that form depend on the pH of the solution. Bjerrum plot:

a. Rxn 0: carbon dioxide dissolution in water

 $CO₂ + H₂O \rightarrow H₂CO₃$

We assume that $f_{CO2} = P_{CO2}$ (in air $P_{CO2} = 10^{-3.5}$ atm), then $K_{\infty} = \frac{a_{H_2CO_3}}{H_2CO_3}$

$$
K_{CO_2} = \frac{n_2 \cos 3}{a_{H_2O} P_{CO_2}}
$$

Rearrange to find a_{H2CO3}:

$$
a_{H_2CO_3} = K_{CO_2}P_{CO_2}
$$

b. Rxn 1: $H_2CO_3 \rightarrow HCO_3 + H^+$

$$
K_1 = \frac{a_{HCO_3^-} a_{H^+}}{a_{H_2CO_3}}
$$

c. Rxn 2: $HCO_3^ \rightarrow$ $CO_3^2^-$ + H^+

$$
K_2 = \frac{a_{CO_3^{2-}}a_{H^+}}{a_{HCO_3^-}}
$$

We usually report K value as pK, where $pK = -\log(K)$. This is similar to H⁺ and pH. At 25° C:

C. Alkalinity and Titrations

How do we actually quantify buffering capacity? The *buffering capacity* of water in the environment is measured as *alkalinity* (equivalents)

Total alkalinity – the equivalent sum of the bases that are titratable with a strong acid; e.g. can accept a proton to form a neutral species

 $\text{Alk}_{\text{tot}} = m_{\text{HCO3-}} + 2 m_{\text{CO32-}} + m_{\text{B(OH)}4} + m_{\text{H3SIO4-}} + m_{\text{organic anions}} + m_{\text{OH}} - m_{\text{H}+}$

$$
Alkalinity\left(\frac{eq}{L}\right) = \sum m_{cations} z_{cations} - \sum m_{conserv,anions} z_{conserv,anions}
$$

Alkalinity is measured by titration:

At end point of weak acid/base cases, $pK_a = pH - log([A-]/[HA])$ (Henderson-Hasselbalch Eqn.) For HA \rightarrow H⁺ + A:

Henderson-Hasselbalch equation can be derived from:

$$
K_a = \frac{a_{A} - a_{H^+}}{a_{HA}}
$$

II. Carbon species as a fraction (α) of total carbon concentration If we define:

$$
\alpha_0 = \frac{[H_2CO_3]}{[H_2CO_3] + [HCO_3^-] + [CO_3^{2-}]} = \frac{[H_2CO_3^1]}{C_T},
$$

\n
$$
\alpha_1 = \frac{[HCO_3^-]}{[H_2CO_3] + [HCO_3^-] + [CO_3^{2-}]} = \frac{[HCO_3^-]}{C_T},
$$

\n
$$
\alpha_2 = \frac{[CO_3^{2-}]}{[H_2CO_3] + [HCO_3^-] + [CO_3^{2-}]} = \frac{[CO_3^{2-}]}{C_T}.
$$

We can derive:

$$
\alpha_0 = \frac{\{H^+\}^2}{\{H^+\}^2 + \{H^+\}K_{a1} + K_{a1}K_{a2}}
$$

$$
\alpha_1 = \frac{\{H^+\}K_{a1}}{\{H^+\}^2 + \{H^+\}K_{a1} + K_{a1}K_{a2}}
$$

$$
\alpha_2 = \frac{K_{a1}K_{a2}}{\{H^+\}^2 + \{H^+\}K_{a1} + K_{a1}K_{a2}}
$$

Procedures of deriving α_0 :

$$
\alpha_0 = \frac{[H_2 C O_3]}{[H_2 C O_3] + [H C O_3^-] + [C O_3^{2-}]}
$$

From $K_1 = \frac{[H^+] [HCO_3^-]}{[H_2CO_3]}$, we can derive $[HCO_3^-] = \frac{K_1 [H_2CO_3]}{[H^+]}$

From $K_2 = \frac{[H^+][CO_3^{2-}]}{[HCO^{-}]}$ $\frac{K_1![CO_3^{2-}]}{[HCO_3^-]}$, we can derive $[CO_3^{2-}] = \frac{K_2[HCO_3]}{[H^+]} = \frac{K_1K_2[H_2CO_3]}{[H^+]^2}$

Substitute the latter two equations into the first one, we can derive the α_0 equation.

III.Example

EXAMPLE 8.1: Hypochlorous acid (HOCl) is commonly used as a disinfectant. If 5 mL of a 10 g/L solution of HOCl is diluted to a total volume of 1 L using distilled water, what is the resulting pH? (Ignore activity corrections.)

Answer: Because activity corrections are being ignored, the problem can be solved on a concentration basis. We have four species (not including H_2O), H^+ , OH⁻, HOCl, OCl⁻, and thus need four equations.

Equilibrium constants:

(1) From Table 8.2: HOCl \rightleftharpoons H⁺ + OCl⁻ $K_a = 10^{-7.60} = \frac{[H^+][OCl^-]}{[HOC]}$ $m+1$

(2)
$$
K_w = [H^+][OH^-] = 10^{-14}
$$

Mass balance:

(3)
$$
OCl_T = [HOC1] + [OC1^-]
$$
, but what is OCl_T ?

We have

$$
5 \text{ mL} \times \frac{1 \text{ L}}{1000 \text{ mL}} \times \frac{10 \text{ g HOC1}}{\text{L}} \times \frac{\text{mol HOC1}}{52.45 \text{ g HOC1}} = 9.5 \times 10^{-4} \text{ mol HOC1 added},
$$

and thus $OCl_T = 9.5 \times 10^{-4}$ M.

Charge balance:

(4) $[H^+] = [OH^-] + [OCl^-]$

Note that the charge balance and proton balance are the same for this problem. We now have four equations and four unknowns. Because we have added an acid to the water, it is reasonable to assume that the pH will be less than 7 and that $[OH^-]$ is negligible. Thus, $[H^+] = [OCl^-]$.

Using Eq. (1) ,

$$
[\text{HOC}]=\frac{[\text{H}^+][\text{OC}^-]}{K_\text{a}}
$$

Substituting into Eq. (3) gives

$$
OCI_T = \frac{[H^+][OCI^-]}{K_a} + [OCI^-] = [OCI^-] \left(1 + \frac{[H^+]}{K_a}\right),
$$

or

$$
[OCI^-] = \frac{OCI_T K_a}{K_a + [H^+]}
$$

Substituting into the simplified charge balance gives

$$
[H^+] = \frac{OCl_T K_a}{K_a + [H^+]}
$$
 or $[H^+]^2 + K_a[H^+] - OCl_T K_a = 0$.

Using the quadratic equation to solve for [H⁺] gives [H⁺] = 4.9 \times 10⁻⁶ (which, as indicated above, is also the concentration of OCl⁻) and $pH = 5.31$. Checking the assumption that [OH]⁻ << [OCl⁻] was correct, we find [OH⁻] = 2.04 × 10⁻⁹ at pH 5.31, and the assumption is valid.

EXAMPLE Determine the alkalinity of solution containing 2×10^{-3} M NaHCO₃, 3 x 10⁻³ M Ca(OH)₂, 1 x 10⁻³ M KOH, 1 x 10⁻³ M H₂SO₄, and 6 x 10⁻³ M HCl.
(1) $[Na^+] + 2[Ca^{2+}] + [K^+] + [H^+] = [HCO_3^-] + 2[CO_3^{2-}] + [OH^-]$ $+2[SO₄²⁻]+ [Cl⁻].$

The sodium, calcium, potassium, sulfate, and chloride concentrations are known, so rearrange to

(2) $[Na^+] + 2[Ca^{2+}] + [K^+] - 2[SO_4^{2-}] - [Cl^-] = [HCO_3^-] + 2[CO_3^{2-}]$
+[OH⁻] -[H⁺].

The right-hand side is [Alk], and thus

(3) $[Alk] = [Na^+] + 2[Ca^{2+}] + [K^+] - 2[SO_4^{2-}] - [Cl^-] = 2 \times 10^{-3}$
 $+ 2(3 \times 10^{-3}) + 1 \times 10^{-3} - 2(1 \times 10^{-3}) - 6 \times 10^{-3} = 1 \times 10^{-3}$ M.

EXAMPLE Calculation of alkalinity and related carbonate species from titration data: A 100-mL water sample with an initial pH of 7.20 required 8.4 mL of 0.0205 N HCl to reach the total alkalinity endpoint (V2). Compute the total alkalinity and determine the equilibrium concentrations of carbonate species in the initial solution. Answer:

Total alkalinity = 8.4 mL \times 0.0205 meq/mL \times 1000 mL/L \div 100 mL $= 1.722$ meq/L,

or

1.722
$$
\text{meq}/L \times 50 \text{ mg } \text{CaCO}_3/\text{meq} = 86.1 \text{ mg/L as } \text{CaCO}_3
$$
.

Because the initial pH is \lt 8.3, there is no carbonate (or phenolphthale in) alkalinity, and all of the total alkalinity is bicarbonate. Thus, $[\text{HCO}_3^-] = 1.722 \times 10^{-3} \text{ M (or } 10^{-2.76} \text{ M}).$ We can compute $[H_2CO_3^*]$ and $[CO_3^{2-}]$ from the given pH, computed $[HCO_3^-]$, and appropriate acid dissociation constant. Because we do not know the ionic strength of the water, we will ignore activity corrections and use the thermodynamic values for K_{a1} and K_{a2} from Table 8.6. Assume the analyses were done at room temperature (20 $^{\circ}$ C). Then $K_{a1} = 10^{-6.38}$ and $K_{a2} = 10^{10.38}$:

$$
[H_2CO_3] = [H^+][HCO_3^-]/K_{a1} = 10^{-7.35} \times 10^{-2.76} \div 10^{-6.38} = 10^{-3.73} \text{M or } [\text{H}_2CO_3]
$$

= 1.86 × 10⁻⁴ M,

and

$$
[CO_3^{2-}] = K_{a2}[HCO_3^-]/[H^+] = 10^{-10.38} \times 10^{-2.76} \div 10^{-7.35} = 10^{-5.79} \text{M or } [CO_3^{2-}]
$$

= 1.62 × 10⁻⁶ M.